Simulating Solute Transport in Porous Media Using Model Reduction Techniques

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Abstract

In this study, we introduce a numerical method to reduce the original transport equations into a reduced form that can replicate the behavior of the original model. The basic idea is to collect an ensemble of data of state variables (say, solute concentration), called snapshots, by running the original model, and then use the proper orthogonal decomposition (POD) techniques (or the Karhunen-Loeve decomposition) to create a set of basis functions that span the snapshot collection. The snapshots can be reconstructed using these basis functions. The solute concentration at any time and location in the domain is expressed as a linear combination of these basis functions, and a Galerkin procedure is applied to the original model to obtain a set of ordinary differential equations for the coefficients in the linear representation. The accuracy and computational efficiency of the reduced model have been demonstrated using several one-dimensional and two-dimensional examples with variable permeability field K_s and sorption coefficient K_d .

Key words: Model reduction, solute transport, proper orthogonal decomposition

1. Introduction

- Accurate predictions of radionuclide transport in general come from pro-
- s cess models, which are defined as detailed flow and contaminant transport
- 4 model that best replicate the available data for a site. They usually are the
- 5 most complex and sophisticated models of flow and transport at a particular

site. A significant technical issue arises when one tries to use those model results in probabilistic systems modeling: it is impractical to directly solve the computationally intensive model for each Monte Carlo realization that would be necessary to properly span the range of uncertainty for every model parameter.

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On contrast to process models, systems models have been proposed to represent complex process models with simplified models that are suitable for Monte Carlo analysis. Systems models incorporate streamlined versions of one or more process models, along with uncertain estimates of the contaminant source term. Systems models integrate knowledge of all of the processes relevant to assessing risk, and therefore are critical to support remedies or decisions to monitor groundwater. Process models synthesize knowledge of one component of the transport path; the systems models integrate knowledge of all aspects of contaminant transport.

The groundwater pathway in the systems model should be an abstraction model that we believe captures the essential features of the groundwater process model, although there is a risk that the simplification process will filter out something important. When the groundwater pathway is simulated in a process model, it is generally not practical to embed the entire model in the systems model - computational limitations intercede. Consequently, we often settle for one or a few simplified models. These simplified models are called abstraction models. Within these simplified or other types of transport modules, parameters that can be varied in a probabilistic analysis include the groundwater velocity, dispersion coefficient, and sorption parameters. In the best case, the process model is used to justify these parameter distributions, but often this is not done in great detail due to the complexity of the process models and the lack of a convenient tool to formally abstract the process model results. The result may be a systems model that does not fully consider, or indeed even contradicts, one or more of the process models. The credibility of the systems model is placed in jeopardy when this happens.

If an efficient and accurate model abstraction procedure can be implemented for groundwater contaminant transport, we largely avoid the issue of justifying the validity of our abstraction model: the original process model is effectively incorporated in the systems model. Alternatively, if a modeling analysis stops short of systems analysis, the reduced model can still be used to explore uncertainties much more efficiently, thereby allowing better sensitivity analyses to be conducted.

To date, methods for proceeding from complex process models to simpli-

fied systems models have not been formally determined or defined. In this study a method is being developed to obtain efficient and accurate reduced models, alternatively referred to as abstraction models. The technique, called the proper orthogonal decomposition (POD), also known as principal component analysis or Karhunen-Loeve expansion, uses the results from a process model as "data" that provides the basis for a reduced model. The procedure, in general terms, calls for the running of the process model, in this case a contaminant transport simulation, forward in time, recording "snapshots" of contaminant concentration. The mathematical technique transforms the results into a set of basis functions that span the behavior of the transport problem throughout the model domain. The reduced model is then constructed by assuming that the solution of the transport problem can be formulated as a linear combination of the basis functions. A Galerkin procedure using the basis functions results in a small set of ordinary differential equations that are solved in time.

There are a variety of applications in various fields of science for which this technique has been applied. For example, researchers in the field of fluid dynamics have used POD methods to discern so-called coherent structures within a turbulent flow field [Berkooz et al., 1993], and to characterize the spatio-temporal patterns in two-phase, fluidized bed reactors [Cizmas et al., 2003]. Due to the need for compact models suitable for integration into process control systems, the techniques have been applied to the modeling of non-linear heat transfer [Park and Cho, 1996], natural convection [Ly and Tran, 1999], and transport of chemicals in chemical vapor deposition (CVD) reactors [Newman, 1996b; Ly and Tran, 1998]. In the field of groundwater hydrology and reservoir engineering, techniques to develop reduced models have been explored [Markovinovic et al., 2002; Hejin et al., 2003; Vermeulen et al., 2004], although these applications focus on fluid flow rather than contaminant transport.

An excellent summary of the underpinning theoretical development of POD is presented by Newman [1996a], and several of the aforementioned references contain detailed descriptions of the implementation of the POD technique. Of greatest relevance in the present study are the CVD model studies [Newman, 1996b; Ly and Tran, 1998]. In these models, fluid flow (a carrier gas) is assumed to be at steady state, and a reacting chemical species is transported in the fluid. Except that the flow is compressible in the CVD case, this is analogous to the problem of chemical transport in groundwater. Thus, although the POD technique has not been developed

for contaminant transport in groundwater, the scope of work required to develop this capability is reasonably constrained due to this previous work.

4 2. Transport Equations

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The governing equation for transport of a single solute in porous media can be written as $[Zyvoloski\ et\ al.,\ 1997]$

$$\frac{\partial(\phi S_l \rho_l C)}{\partial t} = \nabla \cdot [\phi S_l \mathbf{D} \nabla(\rho_l C)] - \nabla \cdot (\mathbf{u}_l \rho_l C) - \rho_r \frac{dC_r}{dt} + q, \tag{1}$$

subject to an initial condition $C(\mathbf{x}, 0) = C_0(\mathbf{x})$ and appropriate boundary conditions. Here ϕ is the porosity, S_l is the liquid saturation, ρ_l is the liquid density, C is the solute concentration, \mathbf{D} is the dispersion tensor, \mathbf{u}_l is the Darcy velocity, C_r represents the adsorption of the solute onto the porous media, $\rho_r \partial C_r / \partial t$ is an equilibrium sorption term, and q is sources or sinks. All parameters in (1) can be space-dependent, but for simplicity the coordinate $\mathbf{x} = (x_1, x_2, x_3)^T$ has been suppressed. The general equilibrium model for adsorption of species onto the porous media is given by $[Polzer\ et\ al., 1992]$:

$$C_r = \frac{\alpha_1 C_l^{\beta}}{1 + \alpha_2 C_l^{\beta}},\tag{2}$$

where α_1 , α_2 , and β are parameters defining different sorption-isotherm models. In this study, the linear isotherm model ($\alpha_1 = K_d$, $\alpha_2 = 0$, and $\beta = 1$) is used: $C_r = K_d C_l$, where K_d is the partition or distribution coefficient.

If we assume that flow is in saturated porous media at the steady state condition, it follows that that ρ_l , \mathbf{D} , and \mathbf{u}_l are independent of time and $S_l \equiv 1$. We may also assume that ρ_l is a constant. In this case, (1) can be simplified as

$$\frac{\partial C}{\partial t} = \frac{\rho_l}{\phi \rho_l + \rho_r K_d} \left\{ \nabla \cdot [\phi \mathbf{D} \nabla(C)] - \nabla \cdot (\mathbf{u}_l C) + q \right\}. \tag{3}$$

This is the full model we deal with in this study. Given appropriate initial and boundary conditions, as well as other parameter fields such as permeability, one can first solve for the velocity field and then the concentration field from (3).

3. Proper Orthogonal Decomposition (POD)

Let $C_i(\mathbf{x}) = C(\mathbf{x}, t_i)$, $i = \overline{1, N_s}$, denote a set of N_s observations (or snapshots) of a state variable (in this case, solute concentration) observed or simulated from the full model run at time t_i , $i = \overline{1, N_s}$. The basic idea of the POD method is to find function $\psi(\mathbf{x})$, called basis functions, that have a structure typical of the members of the ensemble $C_i(\mathbf{x})$. The basis functions are chosen to give the best representation of the ensemble of snapshots, which maximizes

$$\frac{1}{N_s} \sum_{i=1}^{N_s} |(C_i(\mathbf{x}), \psi(\mathbf{x}))|^2 / (\psi(\mathbf{x}), \psi(\mathbf{x})), \tag{4}$$

where (C_i, ψ) is the inner product of the basis function ψ and the concentration field C_i . It can be shown that the basis functions can be expressed as a linear combination of snapshots [Newmana, 1996]:

$$\psi_n(\mathbf{x}) = \sum_{k=1}^{N_s} f_{kn} C_k(\mathbf{x}), \ k = \overline{1, N_s}, \tag{5}$$

where f_{kn} is the k^{th} component of the n^{th} eigenvector of the kernel **K** that is computed from

$$K_{ij} = \frac{1}{N_s} \int_{\Omega} C_i(\mathbf{x}) C_j(\mathbf{x}) d\mathbf{x}.$$
 (6)

By equation (5), each snapshot can be reconstructed exactly using these basis functions as

$$C_k(\mathbf{x}) = \sum_{i=1}^{N_s} f_{ki} \psi_i(\mathbf{x}). \tag{7}$$

To summarize, for a given ensemble of N_s snapshots, one first computes the kernel **K** from (6) and solves the eigenproblem $\mathbf{Kf} = \lambda \mathbf{f}$, which gives N_s eigenvalues λ_n , $i = \overline{1, N_s}$, sorted from largest to smallest, and their corresponding eigenfunctions $f_n(\mathbf{x})$, $i = \overline{1, N_s}$. The basis functions are then computed from (5) and any snapshot can be reconstructed using (7). Since the matrix **K** is real, symmetric and positive semi-definite, all eigenvalues are non-negative. The importance of the n^{th} eigenfunction depends on its relative "energy", characterized by the ratio of the n^{th} eigenvalue to the sum of all eigenvalues (total energy): $\lambda_n / \sum_{i=1}^{N_s} \lambda_i$. In many cases, the first few eigenvalues carry most of the total energy and thus concentration $C(\mathbf{x})$ in (7) can be approximated by truncating the first M terms ($M << N_s$).

To find the spatial distribution of concentration at a time that is not in the ensemble of snapshots, one has to solve for the coefficients in (7). In the following section, a reduced model is introduced, in which ordinary differential equations for these coefficients are derived from the original (partial differential) transport equations.

4. Reduced Model

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Suppose that the full model (3) is solved to obtain a set of snapshots of the concentration distribution $C(\mathbf{x}, t_i)$, $i = \overline{1, N_s}$. Based on the algorithm described in the previous section, one can find a set of basis functions $\psi_i(\mathbf{x})$, $i = \overline{1, N_s}$. Then a Galerkin's method can be employed by seeking an approximation of concentration $C(\mathbf{x}, t)$ as

$$\widehat{C}(\mathbf{x},t) = \sum_{m=1}^{M} a_m(t)\psi_m(\mathbf{x}), \tag{8}$$

where $M < N_s$, and $a_m(t)$ are time-varying coefficients that are independent of spatial locations. The physical mean of (8) is that the concentration field is approximated by a linear combination of some pre-determined, space-dependent "structures" weighted by time-dependent coefficients. Since $\widehat{C}(\mathbf{x},t)$ in (8) is an approximation of the true concentration, replacing $C(\mathbf{x},t)$ in the full model, i.e., transport equation (3), by $\widehat{C}(\mathbf{x},t)$ will in general produce a model error. In the Galerkin's method, this error is forced to be orthogonal to all these basis functions, i.e.,

$$\int_{\Omega} \left[\frac{\partial \widehat{C}}{\partial t} - \frac{\rho_l}{\phi \rho_l + \rho_r K_d} \left\{ \nabla \cdot \left[\phi \mathbf{D} \nabla \widehat{C} \right] - \nabla \cdot (\mathbf{u}_l \widehat{C}) + q \right\} \right] \psi_i(\mathbf{x}) d\mathbf{x} = 0, \quad (9)$$

for $i = \overline{1, M}$. Here the orthogonality of two functions $f(\mathbf{x})$ and $g(\mathbf{x})$ is defined as $\int_{\Omega} f(\mathbf{x})g(\mathbf{x})d\mathbf{x} = 0$. Substituting (8) into (9) and recalling the orthogonality of basis functions yield

$$\frac{da_i(t)}{dt} = \sum_{j=1}^{M} a_j(t) \frac{\rho_l}{\phi \rho_l + \rho_r K_d} \int_{\Omega} \left\{ \nabla \cdot \left[\phi \mathbf{D} \nabla (\psi_j(\mathbf{x})) \right] - \nabla \cdot \left[\mathbf{u}_l \psi_j(\mathbf{x}) \right] \right\} \psi_i(\mathbf{x}) d\mathbf{x}
+ \int_{\Omega} \frac{q(\mathbf{x}, t) \psi_i(\mathbf{x})}{\phi \rho_l + \rho_r K_d} d\mathbf{x},$$
(10)

which can be written in a matrix form as

$$\frac{dA(t)}{dt} = B A(t) + Q(t), \tag{11}$$

where $A = (a_1, a_2, \dots, a_M)^T$, matrix $B = (b_{ij})_{M \times M}$ and vector $Q = (q_i)_{M \times 1}$ are defined as

$$b_{ij} = \frac{\rho_l}{\phi \rho_l + \rho_r K_d} \int_{\Omega} \left\{ \nabla \cdot \left[\phi \mathbf{D} \nabla (\psi_j(\mathbf{x})) \right] - \nabla \cdot \left[\mathbf{u}_l \psi_j(\mathbf{x}) \right] \right\} \psi_i(\mathbf{x}) d\mathbf{x}, \ i, j = \overline{1, M},$$
(12)

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$$q_i(t) = \frac{1}{\phi \rho_l + \rho_r K_d} \int_{\Omega} q(\mathbf{x}, t) \psi_i(\mathbf{x}) d\mathbf{x}, \quad i = \overline{1, M}.$$
 (13)

The initial condition for (11) is derived from the original initial condition $C(\mathbf{x},t) = C_0(\mathbf{x}),$

$$a_i(0) = \int_{\Omega} C(\mathbf{x}, 0) \psi_i(\mathbf{x}) d\mathbf{x}, \quad i = \overline{1, M}.$$
 (14)

Thus the partial differential equation (PDE) has been reduced to a system of M ordinary differential equations (ODE). Provided that M is fairly small, the reduction in computational time should be significant compared to the numerical solution of the PDE in the original model.

5. Illustrative Examples

In this section, several examples are presented to illustrate how the model reduction techniques can significantly reduce computational efforts in solute transport in porous media, while retaining accuracy. This is accomplished by comparing results from the reduced models against those from the full model runs. Note that, in solving the full model numerically, it is quite often that numerical errors may be introduced. To avoid this, in the first two examples, simple one-dimensional transport problems were chosen because analytical solutions for these simple cases are available, which made it easy to assess the accuracy of the reduced model.

5.1. One-dimensional solute transport with linear sorption

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In the first example, we consider a one-dimensional transport problem in a 188 saturated column of 1m in length, uniformly discretized into 200 elements. 189 The hydraulic conductivity is a constant $K_s = 1.0 \ m/day$ for the entire 190 column, and the flow is driven by a hydraulic gradient of dh/dx = 0.001, 191 which produces a uniform flow with Darcy velocity of $1.1813 \times 10^{-8} \ m/s$. Other transport parameters are given as: the dispersivity coefficient α_L 193 0.03333 m, partition coefficient $K_d = 0.1$, water density $\rho_l = 1000.0 kg/m^3$, rock density $\rho_r = 2500 \ kg/m^3$, and porosity $\phi = 0.25$. Under the given initial condition C(x,0) = 0, a fixed concentration C(x,t) = 1.0 at the inlet (x=0), and a zero concentration gradient $\partial C(x,t)/\partial x=0$ at $x=\infty$, this problem can be solved analytically [Batu, 2006]:

$$C(x,t) = \frac{1}{2} \operatorname{erfc} \left[\frac{x - vt/R}{\sqrt{4\alpha_L vt/R}} \right] + \frac{1}{2} \exp\left(\frac{x}{\alpha_L}\right) \operatorname{erfc} \left[\frac{x + vt/R}{\sqrt{4\alpha_L vt/R}} \right], \quad (15)$$

where $R = 1 + \rho_r K_d/(\phi \rho_l)$ is the retardation factor, and *erfc* is the complementary error function. Twenty-five concentration snapshots are computed using (15) at time $t = n\Delta t$, where $\Delta t = 20$ days and $n = \overline{1,25}$. Some selected snapshots are illustrated in Figure 5.

Using these 25 snapshots, the kernel \mathbf{K} is computed from (6), and the eigenvalues and eigenfunctions associated with this kernel are solved from $\mathbf{Kf} = \lambda \mathbf{f}$. These eigenvalues and eigenfunctions depend significantly on the choice of snapshots. It is critical that each snapshot be significantly different from all others. The eigenvalues for this set of snapshots are depicted in Figure 5.1. The figure indicates that the first eigenvalue carries about 91% of the total energy and the first 6 eigenvalues carry more than 99.99% of the total energy. The ratio of the accumulative energy to the total energy is a measure that can be used to determine the number of modes needed to achieve a given accuracy.

The first few basis functions are illustrated in Figure 5.1. It is seen from the figure that the magnitude of the first basis function is much larger than that of other basis functions. In addition, the magnitude decreases as the mode number increases, which makes it possible to approximate the concentration field using only a few terms rather than $N_s(=25)$ terms.

Based on these basis functions, all snapshots can be reconstructed using (7). Comparisons of the true snapshots and reconstructed snapshots indicate that at most 6 basis functions are enough to reconstruct these snapshots with sufficient accuracy. Of course, in general, the number of basis functions required to obtain an accurate solution will depend on the parameters $(K_s, \alpha_L, \text{ and } L)$ in the solute transport problem.

The spatial distribution of solute concentration at any give time can be derived from solving the reduced model. The ordinary differential equation (11) with initial condition (14) was solved using the fourth-order Runge-Kutta method with a time step of $\Delta \tau = 1000 \, s$. In particular, the ensemble of those snapshots can be solved from the reduced model (rather than reconstructed from (7)). Four selected concentration distributions computed from the reduced model, as functions of the number of basis functions included, are illustrated in Figure 5.1. Also plotted in the figure is the true (exact) solution from (15). The figure clearly shows that the accuracy of the estimated concentration distribution depends on the number of basis functions included. For this example, including 10 basis functions is enough to produce very accurate results as compared to the exact solutions.

It is interesting to investigate how well the reduced model can predict the concentration distribution at an elapsed time that is different from those times at which the ensemble of snapshots are taken. Figure 5.1 compares the concentration profiles at time t=75 and 525 days computed from the analytical solution and those solved from the reduced model with 15 basis functions. Note that t=525 days is larger than the maximum time of all snapshots used in constructing the reduced model. The comparison in the figure demonstrates that the reduced model can repreduce the true solution accurately.

5.2. One-dimensional transport with pulse input

In the second example, all transport parameters are the same as in the previous example except that (1) the partition coefficient K_d is zero in this case, and (2) instead of a fixed constant concentration at the inlet (x = 0), a unit pulse input is imposed at the inlet for a duration of 5 days. For this simple

case, analytical solution is also available [Batu, 2006]

$$C(x,t) = \frac{1}{2}erfc\left[\frac{x-vt}{\sqrt{4\alpha_L vt}}\right] + \frac{1}{2}\exp\left(\frac{x}{\alpha_L}\right)erfc\left[\frac{x+vt}{\sqrt{4\alpha_L vt}}\right]$$

$$- \frac{1}{2}erfc\left[\frac{x-v(t-t_0)}{\sqrt{4\alpha_L v(t-t_0)}}\right] - \frac{1}{2}\exp\left(\frac{x}{\alpha_L}\right)erfc\left[\frac{x+v(t-t_0)}{\sqrt{4\alpha_L v(t-t_0)}}\right]$$
(6)

where t_0 is the duration of the pulse input, which starts at zero. Using this equation, 25 snapshots were calculated from t = 10 to 250 days at an increment of 10 days. Several selected snapshots are illustrated in Figure 5.2.

Following the POD method, the kernel K was computed from these 25 snapshots using (6); eigenvalues and eigenfunctions were solved from the eigenproblem $\mathbf{Kf} = \lambda \mathbf{f}$; and then basis functions were computed using (5). The set of eigenvalues as a function of the mode is depicted in Figure 5.2. Unlike the previous case where the first eigenvalue carries about 90% of energy, in this case the first one has only 40% of energy and the first 8 eigenvalues carry about 99.99% energy, which means that more basis functions may be required to approximate the solution. This can also be seen from Figure 5.2, where the relative magnitudes of the first few eigenfunctions are more or less the same, while in the previous example the magnitude of the first eigenfunction is much larger than that of other eigenfunctions. A larger number of required basis functions may be attributed to the fact that the patterns of different snapshots are quite different in this example while in the previous example all snapshots have a very similar pattern.

Figure 5.2 compares four snapshots computed from the analytical solution and those solved from the reduced model with different numbers of basis functions included in the reduced model. It is seen from the figure that the reduced model with 10 basis functions is accurate enough. Figure 5.2 shows the comparison of concentration profiles at time t=75 and 275 days derived from the analytical solution and those from solving the reduced model with 10 basis functions. Again, the solutions from the reduced model are nearly identical to the true solutions.

5.3. Two-dimensional solute transport with heterogeneous permeability and
 sorption coefficients

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In the third example, solute transport is modeled in a two-dimensional rectangular domain of size 1 $m \times 0.5 m$, uniformly discretized into 100×50 square elements. No-flow conditions are prescribed at two lateral boundaries. The hydraulic head is fixed at the left and right boundaries as 10.001m and 10.0m, respectively, with a hydraulic gradient J = 0.001. The porous medium is heterogeneous in both the hydraulic conductivity K_s and the sorption coefficient K_d . It is assumed that the log hydraulic conductivity $Y = \log(K_s)$ has a normal distribution and is second-order stationary following an isotropic, exponential covariance function with a correlation length of 0.3 m. The statistics of the log hydraulic conductivity are given as $\langle Y \rangle = 0$, (i.e., the geometric mean of the saturated hydraulic conductivity $K_G = 1.0 \text{ m/day}$ and $\sigma_Y^2 = 0.693$ (coefficient of variation $CV_{K_s} = 100\%$). Figure 5.2a shows the log hydraulic conductivity field generated using the sgsim code in GSLIB [Deutcsh and Jounel, 1998]. In this case, the velocity field is no longer uniform in the flow domain and has to be solved numerically. The steady-state, saturated flow problem was solved using the Finite-Element Heat- and Mass-Transfer code (FEHM) developed by Zyvoloski et al. [1997]. This velocity field was used as input to the solute transport model.

For the transport problem, it is assumed that the initial concentration is zero in the domain and concentration of $C(\mathbf{x},t)=1.0$ is fixed at the middle of the upstream boundary, $(0.0\ m, 0.25\ m)$. It is also assumed the log partition coefficient is uncorrelated with the log hydraulic conductivity and that $\ln(K_d)$ is also second-order stationary field following an isotropic, exponential covariance function with a correlation length of 0.3 m. The statistics of $\log(K_d)$ are given as $\langle \ln(K_d) \rangle = -1.6094$, (i.e., the geometric mean $K_d^G = 0.2$) and $\sigma_{\ln K_d}^2 = 9.95 \times 10^{-3}$ (coefficient of variation $CV_{K_d} = 10\%$). The spatial distribution of K_d generated using the sgsim code of GSLIB is illustrated in Figure 5.2b. Other transport parameters include the longitudinal dispersivity coefficient $\alpha_L = 0.02\ m$, and the transverse dispersivity coefficient $\alpha_T = 10^{-4}\ m$.

The full transport model was run for 200 days using the FEHM code and the concentration distribution was recorded at t = 10n days, where

 $n = \overline{1,20}$, and these 20 concentration fields were taken as snapshots for the model reduction method. Basis functions were then computed from these snapshots using the POD method, and equation (11) with initial condition (14) were solved numerically.

Figure 5.3 compares the spatial distribution of concentration as contour maps at two elapsed times, $t=100\,$ days and $t=200\,$ days, derived from both the full model run and the reduced model with 5 basis functions. The comparison clearly shows that the results from the reduced model can reproduce the full model results very well, even for a contour level as low as C=0.01.

To compare the computational efficiency of these two methods, the CPU time required for the full model run was recorded, in which the maximum time step was set to 2 days while the actual size of the time step was automatically adjusted during the solution process by the program itself. For the transport problem as described above, the FEHM code takes 23.3 hours. The computation time needed for the reduced model depends on the time step used in the fourth-order Runge-Kutta method, which is fixed at $\Delta \tau = 1000$ seconds in this case. The CPU time for the reduced model run is only 78 seconds. Of course, this run time will increase if a small time step was used or a great number of snapshots and functions is required. However, the comparison of model results from the full model and the reduced model indicates that the time step of $\Delta \tau = 1000$ seconds used in the reduced model is small enough for this problem, and the overall comparison indicates that the number of snapshots and basis functions were sufficient.

For the purpose of comparison, an additional example was simulated, by setting $K_d = 0$ in the previous example. In this case, the plume moves faster than in the third example. As shown in Figure 5.3, comparison of the concentration fields derived from the full model and the reduced model at times t = 50 days and t = 100 days also demonstrates that the reduced model can reproduce the full model results with a sufficient accuracy.

6. Conclusions

In this study, we have demonstrated that the advection-dispersion equation can be cast in reduced model form, and a reduced numerical model can be developed that replicates the behavior of the original model. We have derived, from the original model equations, a method for reducing the transport model, a partial differential equation with unknowns at each numerical grid point, to a small number of ordinary differential equations solved in time. The method consists of running the original model to obtain the snapshots of concentration in the model domain, computing the basis functions for the model using the POD technique, and using these basis functions in a Galerkin procedure to obtain the ordinary differential equations of the reduced model.

The accuracy and computational efficiency of the reduced model have been investigated using several one-dimensional and two-dimensional examples with variable permeability field K_s and sorption coefficient K_d . These examples demonstrate that the reduced model can reproduce the full model results very accurately while the computational time (in terms of the CPU time) required for the reduced model is much less than that required for the full model.

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Figure Captions

- **Figure 1**. Selected snapshots computed from Equ. (15) for the one-dimensional solute transport with linear sorption.
- **Figure 2**. The eigenvalues and their accumulative value as functions of modes for the one-dimensional solute transport with linear sorption.
- **Figure 3**. Selected basis functions for the solute transport with linear sorption.
- Figure 4. Accuracy of computed snapshots using the model reduction method with different numbers of basis functions for the one-dimensional solute transport with linear sorption.
- **Figure 5**. Comparison of exact solution and reduced model solution at time t=75 days and 525 days for the one-dimensional solute transport with linear sorption.
- **Figure 6**. Selected snapshots computed from (16) th eone-dimensional solute transport with pulse input.
- **Figure 7**. The eigenvalues and their accumulative value as functions of modes for the one-dimensional solute transport with pulse input.
- Figure 8. Selected basis functions for the one-dimensional solute transport with pulse input.
- Figure 9. Accuracy of computed snapshots with different numbers of basis functions using the model reduction method for the one-dimensional solute transport with pulse input.
- Figure 10. Comparison of exact solution and reduced model solution at time t = 75 days and 275 days for the one-dimensional solute transport with pulse input.
 - Figure 11. Gaussian random fields for (a) $\log K_s$, and (b) K_d .
- Figure 12. Comparison of true concentration fields and modeled fields using the reduced model at two elapsed times for the case with heterogeneities in both the permeability and the distribution coefficient.
- **Figure 13**. Comparison of true concentration fields and modeled fields using the reduced model for two elapsed times for the case with heterogeneous permeability field but without sorption $(K_d = 0)$.